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井中震源的远场波场特征研究

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摘要 井中震源在逆 VSP、随钻地震和采矿地球物理研究中都有广泛应用。满足“小井孔”(井孔半径远小于特征波长)及“远场”(炮检距大于特征波长)假设时,井中震源的远场波场存在解析解。为了检验解析解在不同情况下的适用性,本文使用最速下降积分计算了不满足上述假设时井中震源远场波场的合成地震记录,即半解析解。模型试验表明,解析解只能在同时满足“小井孔”和“远场”假设时使用;当这两个假设条件不满足时,解析解的振幅和波形相对于半解析解会有明显的偏差。随着假设不满足程度的增加,偏差会逐渐增加,并会逐渐影响走时的准确拾取;这种条件下,采用半解析解才能获得准确的井中震源波场。

关键词 井中震源; 远场波场; 解析解; 最速下降积分; 最速下降法

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Far-field wavefield characteristics of downhole seismic sources

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Abstract Borehole sources, whose scope goes far beyond sources in boreholes, are of extreme importance in research with active seismic sources, including deep seismic sounding, reverse vertical seismic profiling (RVSP), seismic while drilling, mining geophysics, etc. Sources used in these studies are all of cylindrical structures, which is the reason why they are called borehole sources and why their wave fields has unique characteristics. Previous studies on borehole sources are mostly based on analytical solutions obtained when small-borehole assumption (the borehole radius is significantly smaller than the characteristic wave length) and far-field assumption (the offset is greater than the characteristic wave length) are satisfied. It is still an open question whether the analytical solutions are applicable to cases that violate the two assumptions.

This study is based on the synthetic seismograms computed by both analytical solutions and semi-analytical solutions. The analytical solutions used in previous studies are obtained through asymptotic analysis, while the semi-analytical solutions are computed by numerical integration. The semi-analytical solutions are of higher accuracy and therefore regarded as “true solutions”.

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Synthetic seismograms from the analytical solutions are compared to true solutions to validate whether the analytical solutions are applicable to certain cases or not. Accuracy is crucial to the comparison. Yet the high oscillation of solutions in frequency-wavenumber domain brings out a great challenge. We developed a brand-new numerical method called Steepest Descent Integration Method (SDIM). The new method is inspired by the Method of Steepest Descent (SDM) in asymptotic analysis that is specially designed for highly oscillatory integral and is the very method used to obtain the analytical solutions. Replacing approximate integration path and approximate integrand in SDM with accurate ones, SDIM breaks the restraints of small borehole and far field and can compute seismograms at arbitrary offset and arbitrary source frequency with extremely high accuracy efficiently. We calculate the seismograms by both SDIM and SDM for a large offset (1000 m, significantly large compared to borehole radius of 0.1 m) and varied source frequency (0.1~1000 Hz). The assumption of small-borehole is violated in high frequency cases, while far-field assumption fails when the frequency is low. The same experiment is conducted for all three basic borehole sources.

The works presented in the paper can be categorized into two parts, namely the new SDIM and comparison of seismograms. The study of SDIM shows that: (1) The solutions of borehole sources problem in frequency-wavenumber domain are highly oscillatory. The oscillation depends on source frequency and offsets. High frequency sources result in severe oscillation, so as large offsets. (2) The oscillation is attributed to Hankel functions in the solutions whose exponential part account for most of it. Hence, exponential functions are used in the derivation of SDIM instead of Hankel functions, making the work much easier. (3) The only difference between SDIM and SDM is the accuracy of the steepest descent path and the integrand. SDIM uses the accurate path and integrand, while SDM uses approximate ones. In addition, four numerical examples are presented in the paper. Each is designed specifically. They demonstrate that: (1) Results from SDIM are identical to ones from SDM when small-borehole assumption and far-field assumption are satisfied, which supports the validity of SDIM. (2) When small-borehole assumption is violated, the SDM results differ much from the SDIM ones that are considered as true results. It infers that the influence from borehole might not be ignored even for far-field wave field. (3) When far-field assumption fails, the results from SDM are inaccurate as well, which means the absolute value of the offset cannot guarantee far-field. The ratio of the offset to the characteristic wave length matters. (4) The same phenomenon occurs in the wave field of all the three basic borehole sources.

Obtaining accurate far-field seismograms is the key problem of borehole sources research. Yet it is challenging because of highly oscillatory integral involved. By taking advantage of the special form of analytical solutions, we developed a brand-new method for computing highly oscillatory wavenumber integration. It completely avoids the oscillation and results in numerical integration of a fully smooth function, leading to synthetic seismograms with high precision. It also allows us to compute P, S and surface waves separately, reducing their mutual interference. Numerical experiments demonstrate that the results from SDM are considerably different from ones from SDIM, in both amplitude and phase when the small-borehole assumption or the far-field assumption fails. Therefore, the SDM has its constraint and SDIM is a better alternative if accurate wave-field information is needed.

Keywords Downhole seismic sources; Far-field wave field; Analytical solution; Steepest Descent Integration Method; Method of steepest descent

1 引言

井中震源是指具有圆柱形结构的震源,常用于人工震源的地震探测和勘探中,例如深地震测深、地震勘探和采矿地球物理等领域.常见的井中震源包括炸药震源,井下的空气枪、水枪、射孔枪、以及钻头震源等(Chen et al., 1990). 钻井的存在给井中震源问题带来了较为复杂的边界条件,使得井中震源波场体现出与常规震源不同的特性(Lee and Balch, 1982; Meredith et al., 1993).

以震源类型划分,井中震源可以大致分为三种:即径向应力源、轴向应力源和旋转应力源,其他震源可以由这三种震源的组合而成.采矿地球物理中的炸药震源可以简化为径向应力源(Blair, 2007, 2010),随钻地震勘探中的钻头震源可以简化成轴向应力源和旋转应力源的组合(Rector and Hardage, 1992).以研究区域划分,井中震源的波场可以分为井内波场和井外波场.井内波场作为声波测井的主要研究区域,前人已经对其有大量而详实的研究(Tsang and Rader, 1979; Cheng and Toksöz, 1981; Tubman et al., 1984; Cheng, 1994; 沈建国和张海澜, 2000);而井外波场的研究还相对较少(Heelan, 1953; Lee and Balch, 1982; Meredith, 1991; 刘银斌等, 1993a, b; 张钊等, 1995; Blair, 2007).但是随着采矿地球物理和随钻地震的发展,井外波场研究的重要性日益明显(Rector and Marion, 1991; Rector and Hardage, 1992; Haldorsen et al., 1995; Poletto, 2005; Vasconcelos and Snieder, 2008; 陆斌等, 2009; 王鹏等, 2009; 黄伟传等 2010; 吴何珍等, 2010).

井中震源的波场研究方法主要分为数值模拟方法(Cheng, 1994; Dong and Toksöz, 1995)和解析法(Heelan, 1953; Tsang and Rader, 1979; Lee, 1986; Meredith et al., 1993; De Hoop et al., 1994; Blair, 2007)两类.而井外波场问题的计算区域(如炮检距为 1000m)和研究对象(如钻井半径为 0.1m)的尺度相差很大,因此在使用数值模拟方法时,如果要求网格尺寸逼近钻井半径,波场模拟效果较好,但计算量会过大;如果网格尺寸过大则不能很好地研究井中震源的地震波场特征(王鹏等, 2009).所以,解析法成为目前研究井外波场的主要手段.

Heelan(1953)在远场近似下得到了有限长度

的三种基本井中震源的远场解析解. Jordan(1962)和 Abo-Zena(1977)使用不同的方法也获得了井中震源的远场解析解. Lee 和 Balch(1982)考虑了钻井中流体的存在,得到含液井的远场解析解. 刘银斌等(1993a, b)将多个震源的解叠加起来,得到了井中震源阵列的远场解析解. 上述研究均基于两个假设条件:(1)井孔半径远小于特征波长;(2)炮检距大于特征波长.本文中我们分别称之为“小井孔”和“远场”假设. Meredith(1991)和 Blair(2007)使用离散波数法对频率波数域的解进行数值积分,得到了精确的远场波场,并分别应用于井间地震和采矿地球物理的研究. 离散波数法通过增加等间距的虚拟震源,将积分转化为求和,是计算波数域积分的常用方法(Bouchon and Aki, 1977; Bouchon, 1978, 2003).但是波数域积分中极点(pole)和支线(branch cut)所代表的丰富物理意义也随之消失(Lapwood, 1949).为了保留波数域积分的优势,井内波场的研究通常采用实轴积分法(Tsang and Rader, 1979; 沈建国和张海澜, 2000).但是当研究井外波场时,由于炮检距一般远大于钻孔半径,实轴积分的积分函数会出现高频振荡的现象,严重影响数值积分的精度和计算量,不适用于井外波场的计算.

为了克服波数域实轴积分中出现的高频振荡现象,本文提出一种基于最速下降路径的数值积分方法,并通过比较数值积分和解析解之间的误差,探讨远场解析解的适用性条件.

2 井中震源波场解析解

2.1 频率波数域的解析解

假设在各向同性弹性全空间中存在一个无限长的圆柱形空洞(半径为 a),用来模拟钻井或者炮眼.井内的震源通过圆柱形的边界对井外空间产生的影响,通常可以用井壁两侧位移或者应力的连续性来传递.本文中我们考虑井中为真空,井壁为自由边界条件的情况,因此只需考虑应力源对井外空间的作用.

为了便于边界条件的描述,我们采用柱坐标系,并将柱坐标系的 z 轴与圆柱形空洞的对称轴重合(图 1a).在柱坐标系下,边界上的应力可以用 σ_r , $\sigma_{r\theta}$, σ_{rz} 三个分量来表示,它们分别对应于径向、旋转和轴向应力源(图 1b).

柱坐标系下的平衡方程(忽略体力项)为

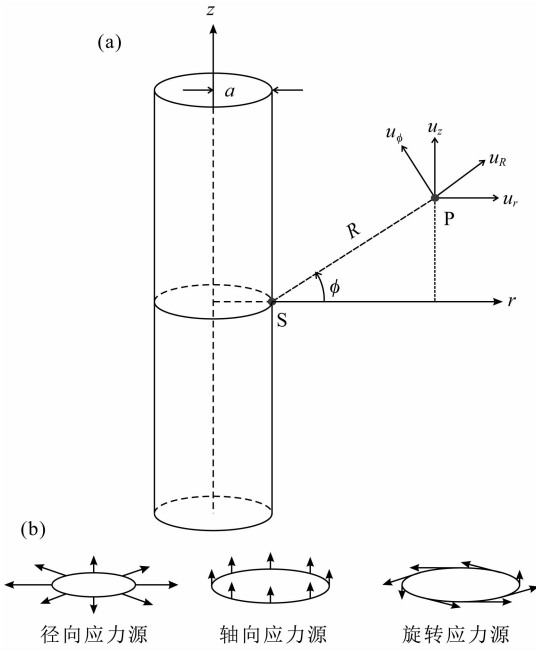


图 1 井中震源示意图

(a)震源结构.震源 S 呈轴对称状分布于半径为 a 的井壁上, 检波器位于 $P(r, z)$ 点. (b)三种基本的井中震源.

Fig. 1 Diagram of downhole seismic sources

(a) Source geometry. The source S is distributed axisymmetrically around the borehole wall of radius a . The receiver is placed at point $P(r, z)$. (b) Three basic types of downhole seismic sources.

$$\begin{aligned} \frac{\partial \sigma_{rr}}{\partial r} + \frac{1}{r} \frac{\partial \sigma_{r\theta}}{\partial \theta} + \frac{\partial \sigma_{rz}}{\partial z} + \frac{\sigma_{rr} - \sigma_{\theta\theta}}{r} &= \rho \frac{\partial^2 u_r}{\partial t^2}, \\ \frac{\partial \sigma_{r\theta}}{\partial r} + \frac{1}{r} \frac{\partial \sigma_{\theta\theta}}{\partial \theta} + \frac{\partial \sigma_{\theta z}}{\partial z} + \frac{2\sigma_{r\theta}}{r} &= \rho \frac{\partial^2 u_\theta}{\partial t^2}, \\ \frac{\partial \sigma_{rz}}{\partial r} + \frac{1}{r} \frac{\partial \sigma_{\theta z}}{\partial \theta} + \frac{\partial \sigma_{zz}}{\partial z} + \frac{\sigma_{rz}}{r} &= \rho \frac{\partial^2 u_z}{\partial t^2}, \end{aligned} \quad (1)$$

各向同性介质的 Hooke 定理是

$$\begin{bmatrix} \sigma_{rr} \\ \sigma_{\theta\theta} \\ \sigma_{zz} \\ \sigma_{\theta z} \\ \sigma_{rz} \\ \sigma_{r\theta} \end{bmatrix} = \begin{bmatrix} \lambda + 2\mu & \lambda & \lambda \\ \lambda & \lambda + 2\mu & \lambda \\ \lambda & \lambda & \lambda + 2\mu \\ & & & \mu \\ & & & & \mu \\ & & & & & \mu \end{bmatrix} \begin{bmatrix} \epsilon_{rr} \\ \epsilon_{\theta\theta} \\ \epsilon_{zz} \\ 2\epsilon_{\theta z} \\ 2\epsilon_{rz} \\ 2\epsilon_{r\theta} \end{bmatrix}, \quad (2)$$

其中 λ, μ 是 Lamé 常数. 几何关系是

$$\begin{aligned} \epsilon_{rr} &= \frac{\partial u_r}{\partial r}, \\ \epsilon_{\theta\theta} &= \frac{u_r}{r} + \frac{1}{r} \frac{\partial u_\theta}{\partial \theta}, \\ \epsilon_{zz} &= \frac{\partial u_z}{\partial z}, \end{aligned}$$

$$\begin{aligned} \epsilon_{r\theta} &= \frac{1}{2} \left(\frac{1}{r} \frac{\partial u_r}{\partial \theta} + \frac{\partial u_\theta}{\partial r} - \frac{u_\theta}{r} \right), \\ \epsilon_{\theta z} &= \frac{1}{2} \left(\frac{\partial u_\theta}{\partial z} + \frac{1}{r} \frac{\partial u_z}{\partial \theta} \right), \\ \epsilon_{rz} &= \frac{1}{2} \left(\frac{\partial u_z}{\partial r} + \frac{\partial u_r}{\partial z} \right), \end{aligned} \quad (3)$$

使用 Helmholtz 分解和极环分解, 我们可以将位移场 $\mathbf{u} = (u_r, u_\theta, u_z)^T$ 分成三个标量场的组合 (Meredith, 1991):

$$\mathbf{u} = \nabla \phi + \nabla \times [\nabla \chi \mathbf{e}_z + \nabla \times (\psi \mathbf{e}_z)], \quad (4)$$

其中 ϕ, ψ, χ 为标量函数, 称为势函数, 分别对应 P 波, SV 波和 SH 波. 将(2)–(5)式代入(1)式, 整理可得,

$$\begin{aligned} \nabla^2 \phi - \frac{\rho}{\lambda + 2\mu} \frac{\partial^2 \phi}{\partial t^2} &= 0, \\ \nabla^2 \psi - \frac{\rho}{\lambda + 2\mu} \frac{\partial^2 \psi}{\partial t^2} &= 0, \\ \nabla^2 \chi - \frac{\rho}{\mu} \frac{\partial^2 \chi}{\partial t^2} &= 0, \end{aligned} \quad (5)$$

其中 $\nabla^2 = \frac{\partial^2}{\partial r^2} + \frac{1}{r} \frac{\partial}{\partial r} + \frac{1}{r^2} \frac{\partial^2}{\partial \theta^2} + \frac{\partial^2}{\partial z^2}$, 是柱坐标系下的 Laplace 算子. 关于 ϕ, ψ, χ 方程的独立性也与各向同性介质中 P, SV, SH 波解耦的现象一致. 震源在该问题中以边界条件的形式表现. 只考虑轴对称问题, 则边界条件可以表示为

$$\begin{aligned} \sigma_{rr} |_{r=a} &= b_1(z, t), \\ \sigma_{r\theta} |_{r=a} &= b_2(z, t), \\ \sigma_{rz} |_{r=a} &= b_3(z, t), \end{aligned} \quad (6)$$

其中 b_1, b_2, b_3 分别对应于径向、旋转和轴向应力源. 将波动方程(5)和边界条件(6)变换到频率波数域, 求解这个边值问题, 可得频率波数域的井中震源解:

$$\begin{bmatrix} \tilde{U}_r \\ \tilde{U}_\theta \\ \tilde{U}_z \end{bmatrix} = \begin{bmatrix} \frac{L_{11}}{d} & 0 & \frac{L_{13}}{d} \\ 0 & L_{22} & 0 \\ \frac{L_{31}}{d} & 0 & \frac{L_{33}}{d} \end{bmatrix} \begin{bmatrix} \tilde{B}_1 \\ \tilde{B}_2 \\ \tilde{B}_3 \end{bmatrix}, \quad (7)$$

其中, $\tilde{U}_r, \tilde{U}_\theta, \tilde{U}_z, \tilde{B}_1, \tilde{B}_2, \tilde{B}_3$ 分别是位移和边界条件在频率波数域中的表示, 满足

$$\begin{aligned} \tilde{U}_j(r, h, \omega) &= \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} u_j(r, z, t) \exp(-ihz + i\omega t) dz dt, j = r, \theta, z \\ \tilde{B}_j(r, h, \omega) &= \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} b_j(r, z, t) \exp(-ihz + i\omega t) dz dt, j = 1, 2, 3 \end{aligned} \quad (8)$$

而 d, L_{ij} 分别为

$$\begin{aligned}
d &= (\rho\omega^2 - 2\mu h^2)^2 H_0^{(1)}(q_1 a) H_1^{(1)}(q_2 a) + 4\mu^2 h^2 q_1 q_2 H_1^{(1)}(q_1 a) H_0^{(1)}(q_2 a) - 2\mu\rho\omega^2 q_1 \frac{1}{a} H_1^{(1)}(q_1 a) H_1^{(1)}(q_2 a), \\
L_{11} &= q_1 [(\rho\omega^2 - 2\mu h^2) H_1^{(1)}(q_2 a) H_1^{(1)}(q_1 r) + 2\mu h^2 H_1^{(1)}(q_1 a) H_1^{(1)}(q_2 r)], \\
L_{13} &= ih \begin{bmatrix} -2\mu q_1 \left(q_2 H_0^{(1)}(q_2 a) - \frac{1}{a} H_1^{(1)}(q_2 a) \right) H_1^{(1)}(q_1 r) \\ + \left[(\rho\omega^2 - 2\mu h^2) H_0^{(1)}(q_1 a) - 2\mu q_1 \frac{1}{a} H_1^{(1)}(q_1 a) \right] H_1^{(1)}(q_2 r) \end{bmatrix}, \\
L_{22} &= -\frac{H_1^{(1)}(q_3 r)}{\mu q_3 H_2^{(1)}(q_3 a)}, \\
L_{31} &= -ih [(\rho\omega^2 - 2\mu h^2) H_1^{(1)}(q_2 a) H_0^{(1)}(q_1 r) - 2\mu q_1 q_2 H_1^{(1)}(q_1 a) H_0^{(1)}(q_2 r)], \\
L_{33} &= -2\mu h^2 \left(q_2 H_0^{(1)}(q_2 a) - \frac{1}{a} H_1^{(1)}(q_2 a) \right) H_0^{(1)}(q_1 r) - \left[(\rho\omega^2 - 2\mu h^2) H_0^{(1)}(q_1 a) - 2\mu q_1 \frac{1}{a} H_1^{(1)}(q_1 a) \right] q_2 H_0^{(1)}(q_2 r).
\end{aligned} \tag{9}$$

其中 ω 是频率, h 是 z 方向波数, q_j ($j=1, 2, 3$) 分别是 P, SV 和 SH 波的 r 方向波数, r, z 是检波器的位置, $H_n^{(1)}(z)$ 是 n 阶 Hankel 函数, 代表 n 阶的柱面波. 获得物理空间域的解需要对频率波数解进行如下反变换:

$$\mathbf{u}(r, z, t) = \left(\frac{1}{2\pi} \right)^2 \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} \tilde{\mathbf{U}}_0(r, h, \omega) \exp[i(hz - \omega t)] dh d\omega, \tag{10}$$

其中 $\tilde{\mathbf{U}}_0 = [\tilde{U}_r, \tilde{U}_\theta, \tilde{U}_z]^T$.

前人关于井中震源的解析解研究中, 尽管采用的具体方法不尽相同, 但最终都会得到与(9)式等价的解(Heelan, 1953; Abo-Zena, 1977; Lee and Balch, 1982; Meredith, 1991; 刘银斌等, 1993a, b; Blair, 2007). 而对(10)式中双重积分的计算方法则主要分为两类. 一类是在远场的假设下, 解析地计算两个积分, 得到远场解析解(Heelan, 1953; Abo-Zena, 1977; Lee and Balch, 1982); 另一类是采用数值积分法计算, 得到半解析解(Meredith, 1991; Blair, 2007).

2.2 近似条件下的远场解析解

建立远场解析解需要两个基本假设条件. 首先是“小井孔”假设, 即钻井半径远小于特征波长. 这时 $qa = 2\pi \frac{a}{\lambda} \ll 1$, 故 Hankel 函数有如下近似表达式:

$$H_n^{(1)}(qa) \approx \begin{cases} i \frac{2}{\pi} \left[\ln\left(\frac{qa}{2}\right) + \gamma \right] & \text{if } n = 0, \\ -i \frac{\Gamma(n)}{\pi} \left(\frac{2}{qa}\right)^n & \text{if } n > 0, \end{cases} \tag{11}$$

其中 γ 是 Euler 常数, $\Gamma(n)$ 是 gamma 函数. 并且可以进一步推知,

$$\frac{H_n^{(1)}(qa)}{H_{n+1}^{(1)}(qa)} \approx 0, \tag{12}$$

所以(9)式中的所有关于 $H_n^{(1)}(qa)$ 的项都可以得到简化, 只保留了 $H_n^{(1)}(qr)$.

当 r 大于特征波长, 即“远场”假设成立时, 可以采用最速下降法(Aki and Richards, 2002)计算关于 $H_n^{(1)}(qr)$ 的波数域积分, 从而得到频率域的解析解:

$$\begin{aligned}
\bar{u}_r(r, z, \omega) &= \frac{ia^2}{4\mu R} \cos\phi \left[\left(1 - \frac{2\mu}{\lambda + 2\mu} \sin^2\phi \right) \frac{\omega}{\alpha} e^{i\omega \frac{R}{\alpha}} + 2\sin^2\phi \frac{\omega}{\beta} e^{i\omega \frac{R}{\beta}} \right] \bar{b}_1 - \frac{a}{2\mu R} \left[\frac{\mu}{\lambda + 2\mu} e^{i\omega \frac{R}{\alpha}} - e^{i\omega \frac{R}{\beta}} \right] \sin\phi \cos\phi \bar{b}_3, \\
\bar{u}_z(r, z, \omega) &= \frac{ia^2}{4\mu R} \sin\phi \left[\left(1 - \frac{2\mu}{\lambda + 2\mu} \sin^2\phi \right) \frac{\omega}{\alpha} e^{i\omega \frac{R}{\alpha}} - 2\cos^2\phi \frac{\omega}{\beta} e^{i\omega \frac{R}{\beta}} \right] \bar{b}_1 - \frac{a}{2\mu R} \left[\frac{\mu}{\lambda + 2\mu} \sin^2\phi e^{i\omega \frac{R}{\alpha}} + \cos^2\phi e^{i\omega \frac{R}{\beta}} \right] \bar{b}_3, \\
\bar{u}_\theta(r, z, \omega) &= \frac{ia^2}{4\mu R} \frac{\omega}{\beta} \cos\phi \bar{b}_2,
\end{aligned} \tag{13}$$

其中 $\bar{u}_r, \bar{u}_\theta, \bar{u}_z, \bar{b}_1, \bar{b}_2, \bar{b}_3$ 分别是频率域内的位移和边界条件, α, β 分别是 P 波和 SV 波的速度, ϕ 是射线方向和柱坐标系中 r 方向的夹角, $R = \sqrt{r^2 + z^2}$.

假设三种应力源的形式都是 $\delta(z)G(t)$, 其中 $G(t)$ 是震源时间函数. 那么对(13)式作反 Fourier 变换, 可以得到这三类震源在物理空间域的解.

(1) 径向应力源:

$$\begin{aligned} u_r(r, z, t) &= -\frac{a^2}{4\mu R} \cos\phi \left[\left(1 - \frac{2\mu}{\lambda + 2\mu} \sin^2\phi\right) \frac{1}{\alpha} G' \left(t - \frac{R}{\alpha}\right) + 2\sin^2\phi \frac{1}{\beta} G' \left(t - \frac{R}{\beta}\right) \right], \\ u_z(r, z, t) &= -\frac{a^2}{4\mu R} \sin\phi \left[\left(1 - \frac{2\mu}{\lambda + 2\mu} \sin^2\phi\right) \frac{1}{\alpha} G' \left(t - \frac{R}{\alpha}\right) - 2\cos^2\phi \frac{1}{\beta} G' \left(t - \frac{R}{\beta}\right) \right]. \end{aligned} \quad (14)$$

(2) 轴向应力源:

$$\begin{aligned} u_r(r, z, t) &= -\frac{a}{2\mu R} \left[\frac{\mu}{\lambda + 2\mu} \sin\phi \cos\phi G(t - R/\alpha) - \cos\phi \sin\phi G(t - R/\beta) \right], \\ u_z(r, z, t) &= -\frac{a}{2\mu R} \left[\frac{\mu}{\lambda + 2\mu} \sin^2\phi G(t - R/\alpha) + \cos^2\phi G(t - R/\beta) \right]. \end{aligned} \quad (15)$$

(3) 旋转应力源:

$$u_\theta(r, z, t) = -\frac{a^2}{4\mu R\beta} \cos\phi G' \left(t - \frac{R}{\beta}\right). \quad (16)$$

其中 $G'(t)$ 为震源时间函数的导数, 是径向和旋转应力源产生的远场波形, 而轴向压力源产生的远场波形则正比于震源时间函数 $G(t)$.

在径向和轴向应力源的位移表示为 P 波和 SV 波的组合. 为了分离 P 和 SV 的贡献, 我们引入坐标系 (R, ϕ, θ) , 其中

$$r = R\cos\phi, z = R\sin\phi, \quad (17)$$

那么 R, ϕ 方向的位移可以通过 r, z 方向的位移旋转得到

$$u_R = u_r \cos\phi + u_z \sin\phi, u_\phi = -u_r \sin\phi + u_z \cos\phi, \quad (18)$$

在新的坐标系下, 径向和轴向应力源的位移分别为

$$u_R = -\frac{a^2}{4\mu\alpha R} \left(1 - \frac{2\mu}{\lambda + 2\mu} \sin^2\phi\right) G' \left(t - \frac{R}{\alpha}\right), \quad (19)$$

$$u_\phi = \frac{a^2}{2\mu\beta R} \sin\phi \cos\phi G' \left(t - \frac{R}{\beta}\right),$$

和

$$u_R = -\frac{a}{2(\lambda + 2\mu)R} \sin\phi G(t - R/\alpha), \quad (20)$$

$$u_\phi = -\frac{a}{2\mu R} \cos\phi G(t - R/\beta).$$

考虑到本文中 ϕ 的定义与前人使用的 ϕ_L 的关系 $\phi = \phi_L - \pi/2$, 上述公式与前人的结果一致 (Heelan, 1953; Lee and Balch, 1982). 从 (19)、(20) 式中可以看出, 两种震源激发的 SV 波能量都比 P 波能量大, 并且能量差会随着两种波速差的增大而增大 (图 2).

3 井中震源波场半解析解

为了计算最速下降路径积分, 我们首先求解最速下降路径的解析表达式, 然后采用数值积分方法, 沿着最速下降路径计算频率波数域解的积分.

理想情况下, 计算 $F(h)$ 的最速下降积分路径

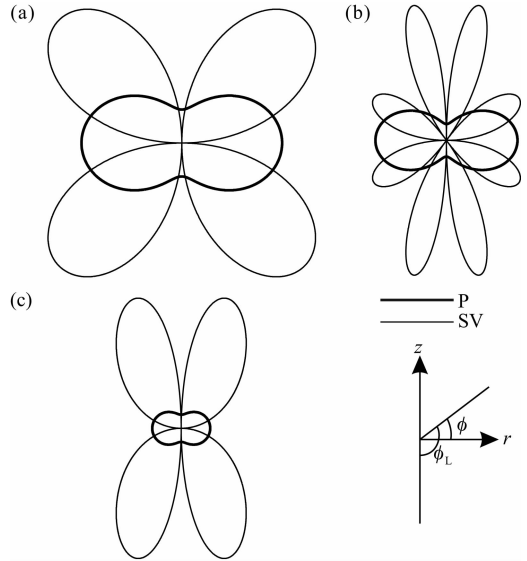


图 2 井中震源的远场辐射图样

(修改自 Lee and Balch, 1982; 参数相同)

震源分别是 (a) 施加在裸眼井井壁上的径向压力源, (b) 施加在充液井井壁上的径向压力源, 和 (c) 位于充液井对称轴上的单极声波源.

Fig. 2 Far-field radiation patterns for downhole seismic sources (modified after Lee and Balch, 1982; same parameters are used)

The sources are respectively (a) a radial stress source applied on an empty borehole, (b) a radial stress source applied on a fluid-filled borehole, and (c) a monopole acoustic source placed at the center of a fluid-filled borehole.

应遵循如下步骤:

$$\int_{-\infty}^{+\infty} F(h) dh = \int_C \exp(\ln F(h)) dh, \quad (21)$$

其中积分函数 $\exp(\ln F(h))$ 的振荡是由 $\ln F(h)$ 虚部的变化导致的. 而沿 $\ln F(h)$ 的虚部等值线进行积分时, $\ln F(h)$ 虚部保持不变, 积分函数的振荡性就会完全消失. 从复平面上的大多数点出发沿虚部等值线积分, 两个方向的积分函数分别是指数增加和指数减少. 但是对于某些特殊的点来说, 两个方向的积分函数都以指数形式减少. 我们称这些点为鞍点 (saddle point), 记为 h_s . 它们满足如下条件:

$$\left. \frac{d}{dh} \ln F(h) \right|_{h=h_s} = 0, \quad \left. \frac{d^2}{dh^2} \ln F(h) \right|_{h=h_s} \neq 0, \quad (22)$$

从鞍点出发的所有方向中,沿虚部等值线两个方向,积分函数的减小是最快的.因此,这条经过鞍点的 $\ln F(h)$ 虚部等值线被称为最速下降路径 (Steepest Descent Path, 简称 SDP). SDP 上任意一点 h 都满足

$$\ln F(h) = \ln F(h_s) - X^2, \quad (23)$$

其中 X 是任意正实数. 已知 h_s 就可利用上式计算出整条最速下降路径.

实际情况下遇到的函数往往形式较为复杂,解析地计算 $F(h)$ 的最速下降路径比较困难. 如果积分函数可以写成 $G(h)F(h)$ 的形式, 其中 $F(h)$ 可以

$$\begin{aligned} \tilde{U}_r &= \tilde{B}_1 \frac{\hat{L}_{111}}{\hat{d}} \exp[iq_1(r-a)] + \tilde{B}_1 \frac{\hat{L}_{112}}{\hat{d}} \exp[iq_2(r-a)] + \tilde{B}_3 \frac{\hat{L}_{131}}{\hat{d}} \exp[iq_1(r-a)] + \tilde{B}_3 \frac{\hat{L}_{132}}{\hat{d}} \exp[iq_2(r-a)], \\ \tilde{U}_\theta &= \tilde{B}_2 \hat{L}_{22} \exp[iq_3(r-a)], \\ \tilde{U}_z &= \tilde{B}_1 \frac{\hat{L}_{311}}{\hat{d}} \exp[iq_1(r-a)] + \tilde{B}_1 \frac{\hat{L}_{312}}{\hat{d}} \exp[iq_2(r-a)] + \tilde{B}_3 \frac{\hat{L}_{331}}{\hat{d}} \exp[iq_1(r-a)] + \tilde{B}_3 \frac{\hat{L}_{332}}{\hat{d}} \exp[iq_2(r-a)], \end{aligned} \quad (25)$$

其中 $\hat{d}, \hat{L}_{11j}, \hat{L}_{13j}, \hat{L}_{31j}, \hat{L}_{33j}, j = 1, 2$ 可以通过将(9)式中的 $H_n^{(1)}(z)$ 替换成 $\hat{H}_n^{(1)}(z)$ 得到, $\hat{L}_{111}, \hat{L}_{112}$ 代表 \hat{L}_{11} 中 $H_n^{(1)}(q_1 r), H_n^{(1)}(q_2 r)$ 的系数. 从(25)式可以看到 $\exp[iq_j(r-a)]$ 项对频率波数域解的振荡性贡献最大. 考虑到波数域积分函数为 $\tilde{U}_0(h, \omega) \exp(iz)$, 我们将 $F(h)$ 设为 $\exp[iq(r-a) + hz]$.

记

$$f(h) = \ln F(h) = i[q(r-a) + hz], \quad (26)$$

其中

$$q = \left(\frac{\omega^2}{c^2} - h^2 \right)^{\frac{1}{2}}, \text{ 取 } \text{Im} q > 0 \cup (\text{Im} q = 0 \cap \omega \text{Re} q > 0), \quad (27)$$

其中 $c = \alpha, \beta$. 与近似条件下的解析解中使用的 $f(h) = i(qr + hz)$ 不同, 我们在最速下降路径的计算中考虑了井孔半径 a 的影响. 令 $f'(h) = 0$, 可以得到

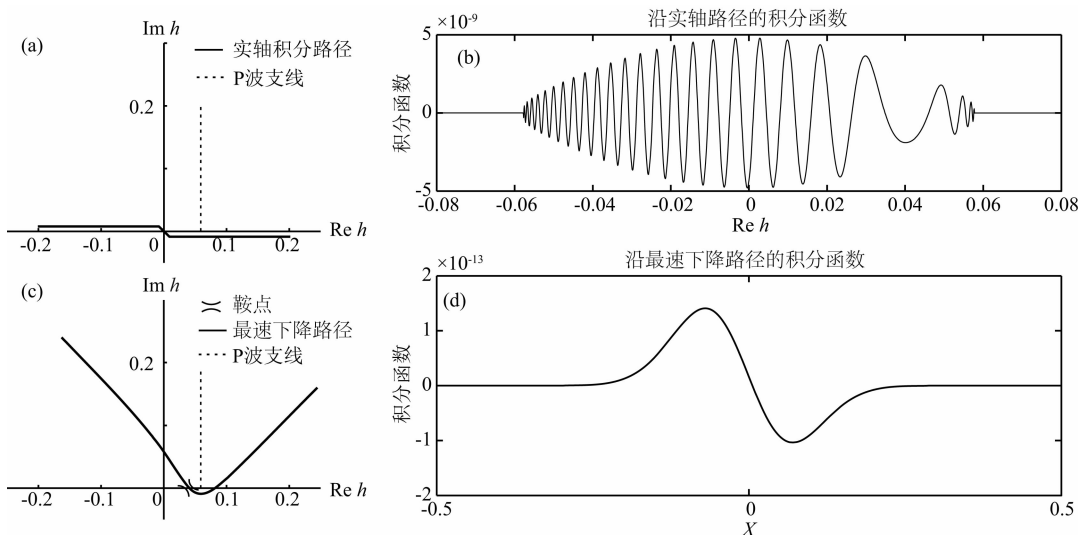


图 3 最速下降积分和实轴积分的对比

图(a)(c)分别是实轴积分和最速下降积分在复波数平面内的积分路径,图(b)(d)则是两个路径上的积分函数. 波数 h 的单位是 m^{-1} .

Fig. 3 Comparison of Steepest Descent Integration Method and Real-axis Integration Method

Panels (a) (c) show two different integration paths for Real-axis Integration Method and Steepest Descent Integration Method in complex wavenumber plane, while (b) (d) are the corresponding integrands along the paths.

$$h_s = \frac{\omega}{c} \frac{z}{R}, q_s = \frac{\omega}{c} \frac{r-a}{R}, f(h_s) = i \frac{\omega}{c} R, \quad (28)$$

其中 $R = \sqrt{(r-a)^2 + z^2}$. 将(26)式代入 $f(h) = f(h_s) - X^2$, 整理可得,

$$R^2 h^2 + [2i(f(h_s) - X^2)z]h - \left[\frac{\omega^2 (r-a)^2}{c^2} + (f(h_s) - X^2)^2 \right] = 0, \quad (29)$$

方程解为

$$h_{1,2} = \frac{-2i[f(h_s) - X^2]z \pm \sqrt{\Delta}}{2R^2}, \quad (30)$$

其中 Δ 是(29)式的判别式. 当 $X = 0$ 时,

$$h_{1,2} = -i \frac{f(h_s)z}{R^2} = \frac{\omega z}{cR} = h_s, \quad (31)$$

当 X 从 0 变化到 $+\infty$ 时, $h_{1,2}$ 形成两条从 h_s 一直到无穷大的曲线, 这两条曲线合并一起就是最速下降路径.

$\sqrt{\Delta}$ 有两个解, 如果设 h_1 是 $\text{Re} \sqrt{\Delta} > 0$ 的那个解, 那么

$$\text{Re} h_1 = \text{Re} \left[\frac{-2i[f(h_s) - X^2]z}{2R^2} \right] + \text{Re} \sqrt{\Delta} = \text{Re} \left[-i \frac{f(h_s)}{R^2} \right] + \text{Re} \sqrt{\Delta} > \text{Re} h_s, \quad (32)$$

因此, h_1 所代表的分支在鞍点 h_s 的右侧, 我们称之为 SDP 的右支, 记为 C_1 ; 同理 h_2 是 SDP 的左支, 记为 C_2 . $C_{1,2}$ 的正方向为 X^2 增加的方向, 所以最速下降路径 $C_{\text{SDP}} = -C_2 + C_1$. 结合复变函数中的 Jordan 引理, 可以将实轴上的积分完全变成沿 SDP 的积分, 如下式所示:

$$\begin{aligned} \int_{-\infty}^{+\infty} \mathbf{U}_0(h, \omega) e^{ihz} dh &= \int_{C_{\text{SDP}}} \mathbf{G}(h, \omega) e^{i[q(r-a) + hz]} dh \\ &= \int_{-C_2} \mathbf{G}(h_2, \omega) e^{i[q(r-a) + h_2 z]} dh + \int_{C_1} \mathbf{G}(h_1, \omega) e^{i[q(r-a) + h_1 z]} dh \\ &= - \int_0^{+\infty} \mathbf{G}(h_2, \omega) e^{-X^2} h'_2(X) dX + \int_0^{+\infty} \mathbf{G}(h_1, \omega) e^{-X^2} h'_1(X) dX \\ &= \int_0^{+\infty} [\mathbf{G}(h_1, \omega) h'_1(X) - \mathbf{G}(h_2, \omega) h'_2(X)] e^{-X^2} dX, \end{aligned} \quad (33)$$

由于积分函数以 $\exp(-X^2)$ 衰减, 根据计算精度要求, 可以很容易地确定上式的积分截断上限. 同时, 由于积分函数非常光滑, 较低的采样率也可以得到相当高的精度.

4 近似条件下远场解析解的适用性特征

4.1 最速下降积分法试验

我们首先验证最速下降积分法在计算远场波场时的适用性, 采用如下参数模型: 井孔半径为 0.1 m, 检波器放置在距离钻井 1000 m 处; 井外地层为常见的 Pierre 页岩, 其 P 波波速为 $2074 \text{ m} \cdot \text{s}^{-1}$, S 波为 $869 \text{ m} \cdot \text{s}^{-1}$, 密度为 $2250 \text{ kg} \cdot \text{m}^{-3}$ (Meredith, 1991); 震源为径向应力震源, 震源时间函数是主频为 30 Hz 的 Ricker 子波(图 4a). 该主频的震源在这种地层中的特征波长 λ_m 约为 346 m, 既远大于井孔半径 2 m, 又远小于 1000 m, 同时满足“小井孔”和“远场”假设, 近似条件下的解析解可以准确地得到位移解.

从径向应力源的远场解析解(14)式可知, 远场

的波形正比于 Ricker 子波的导数(图 4b). 图 4c 是地震记录的径向分量, 其中实线代表半解析解, 虚线代表解析解, 倒三角形标注了 P 波的理论到时. 数值试验结果表明, 用最速下降路径积分得到的数值解与近似条件下的解析解吻合很好(图 4c). 由于在本例条件下, 解析解对位移刻画较为准确, 所以这一结果进一步验证了最速下降积分法的正确性.

4.2 径向应力震源

4.2.1 高频情况

径向应力震源通常用来模拟炮眼中的爆炸震源, 典型的炮眼半径约为 0.1 m. 为了避免震源激发时地面的剧烈振荡破坏检波器和近井的井壁面波的影响, 检波器一般放置在离井一段距离之外, 放置在 $r=1000 \text{ m}$, $z=0 \text{ m}$ 处, 其他参数也同 4.1 节.

我们首先检验“小井孔”假设失效时的情况. 在井孔绝对半径不变的情况下, 通过减小特征波长同样可以使“小井孔”假设失效. 在地层速度不变的前提下, 提高震源的主频可以减小特征波长. 计算结果如图 5 所示. 该模型中, “小井孔”的条件为特征波长 $\lambda \gg a = 0.1 \text{ m}$, 图 5a 中可以看出, 该假设条件完全满

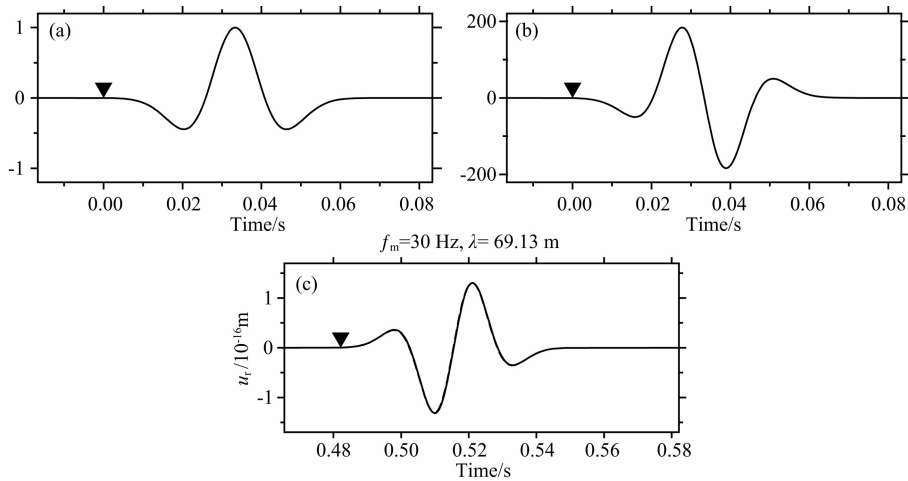


图 4 径向应力源激发波场的半解析解和解析解对比

(a) 震源时间函数(Ricker子波);(b)解析解远场波形(Ricker子波的导数);(c)实线表示最速下降积分得到 u_r 的半解析解, 虚线为解析解. 倒三角形为P波的到时.

Fig. 4 Comparison of the semi-analytical solution and the analytical solution

of wave field excited by a radial stress source

(a) Source time function (Ricker wavelet); (b) Far-field waveform of analytical solution (derivative of Ricker wavelet); (c) Solid line denotes the semi-analytical solution of u_r obtained by Steepest Descent Integration Method and dashed line is the analytical solution. The inverted triangle denotes the theoretical arrival time of P waves.

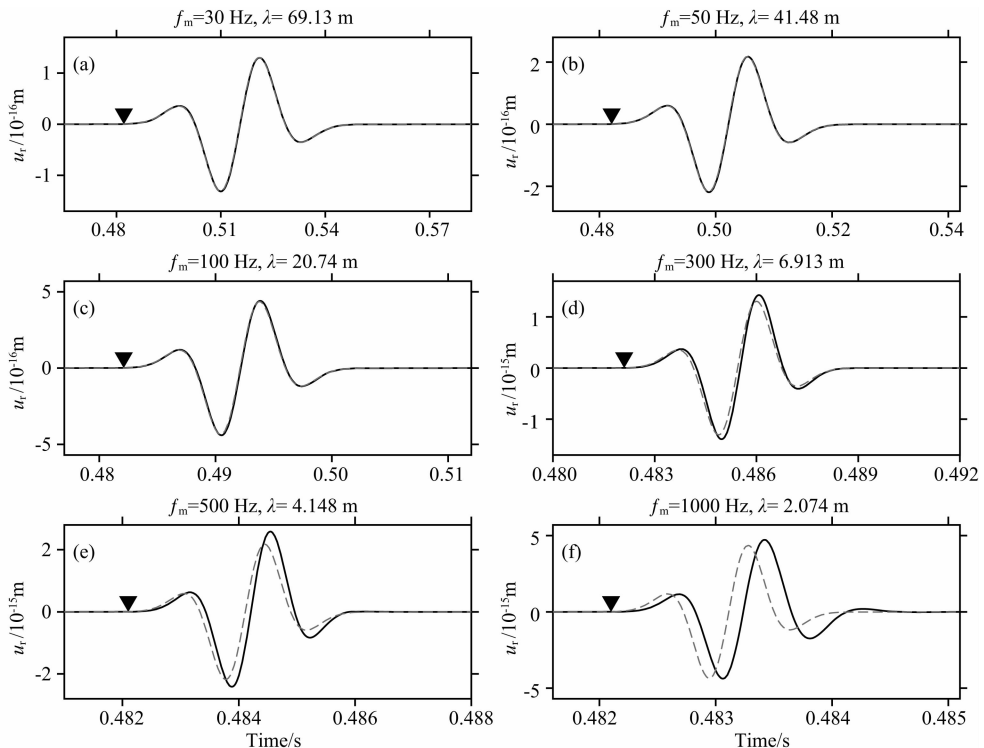


图 5 高频情况下径向应力源激发波场的半解析解(实线)和近似条件下的解析解(虚线)对比

图(a—f)分别是Ricker子波主频为30, 50, 100, 300, 500, 1000 Hz时,位于 $r=1000$ m, $z=0$ m处的检波器记录的径向分量. 倒三角形为P波理论到时.

Fig. 5 Comparison of the semi-analytical solution (solid line) and the analytical solution (dashed line)

of wave field excited by a radial stress source for high frequency

Panels (a—f) are radial components of seismograms recorded by the receiver placed at $r=1000$ m, $z=0$ m with the peak frequency of Ricker wavelet being 30, 50, 100, 300, 500, 1000 Hz respectively. The inverted triangle denotes the theoretical arrival time of P waves.

足时,两者之间差异很小.当震源主频逐渐增加时,子波波长逐渐减小,图 5b 和图 5c 的假设条件接近失效,两者之间的差异增加,但波形差异仍然较小.当子波波长进一步减小时,“小井孔”假设失效(图 5d—5f),虽然数值解和近似条件下的解析解的 P 波到时一直保持在理论到时 0.4821 s 处,但两者在振幅和相位方面的差异开始凸显.其中,半解析解的振幅略高于解析解的振幅,最大振动的到时相对滞后.

解析解将波场对圆柱形井壁这一特殊的边界条件的复杂响应简化为一个尺度因子 a^2 (见式(14)).满足“小井孔”假设时,由于井孔半径远小于特征波长,钻井对远场波场影响不大,解析解的简化较为合理.而当“小井孔”假设失效时,井孔对波场的作用开始体现,即使在远场的地震记录上也会有响应.

除了振幅和相位相对差异外,两者的绝对振幅都会随着频率的变大而迅速地增加.这是因为两者的位移都大致正比于 Ricker 子波的导数.

4.2.2 低频情况

当震源频率变低时,特征波长会随之增加.虽然这时“小井孔”假设不会受到影响,但是有可能会导

致“远场”假设失效.低频情况下的计算结果如图 6 所示.该模型中“远场”假设的条件是 $\lambda \ll r = 1000$ m.从图 6a 中可以看出,该假设完全满足,两者差异很小.但是当震源主频逐渐到 5 Hz 时,“远场”假设接近失效,两者差异已经开始体现(图 6b).随着震源主频的进一步增加,两者之间的差异越来越明显(图 6c—6f).在低频情况下,半解析解和解析解同样出现了振幅和相位上的偏差.数值解的振幅大于解析解,并且比高频情况下更加明显(图 6d—6f).与高频情况不同的是,低频情况下数值解的波形与解析解也有比较明显的差别,逐渐从 Ricker 子波的导数变成了 Ricker 子波.

一般情况下,远场是指炮检距的尺度远大于震源尺度;而在井中震源问题中,“远场”假设比较的是炮检距的尺度和特征波长的尺度.所以虽然本例中炮检距(1000 m)远大于震源的尺度(0.1 m),但是并不一定满足“远场”假设.如图 6c—6f 中,特征波长大于或等于炮检距时,解析解与数值解出现明显的差异.

4.3 轴向应力震源

轴向应力震源会用来模拟一些附着在井壁上的

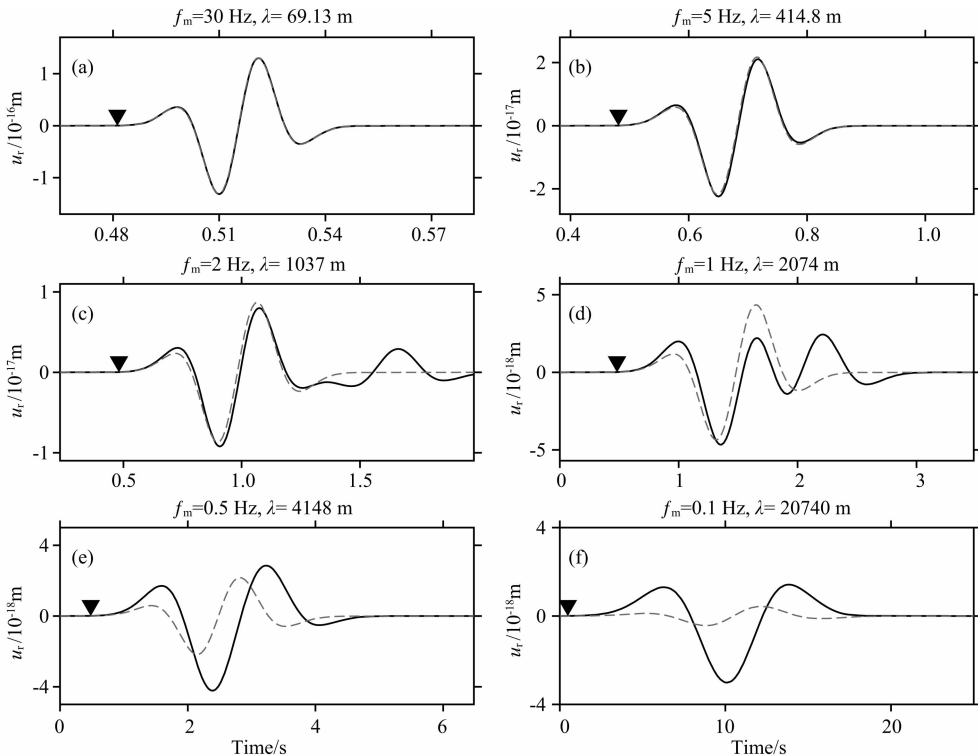


图 6 低频情况下径向应力源激发波场的半解析解(实线)和解析解(虚线)的对比
图(a—f)分别是 Ricker 子波主频为 30, 5, 2, 1, 0.5, 0.1 Hz 时检波器记录的径向分量.倒三角形为 P 波理论到时.

Fig. 6 Comparison of the semi-analytical solution (solid line) and the analytical solution (dashed line) of wave field excited by a radial stress source for low frequency

Panels (a—f) are radial components of seismograms recorded by the receiver placed at $r = 1000$ m, $z = 0$ m with the peak frequency of Ricker wavelet being 30, 5, 2, 1, 0.5, 0.1 Hz respectively. The inverted triangle denotes the theoretical arrival time of P waves.

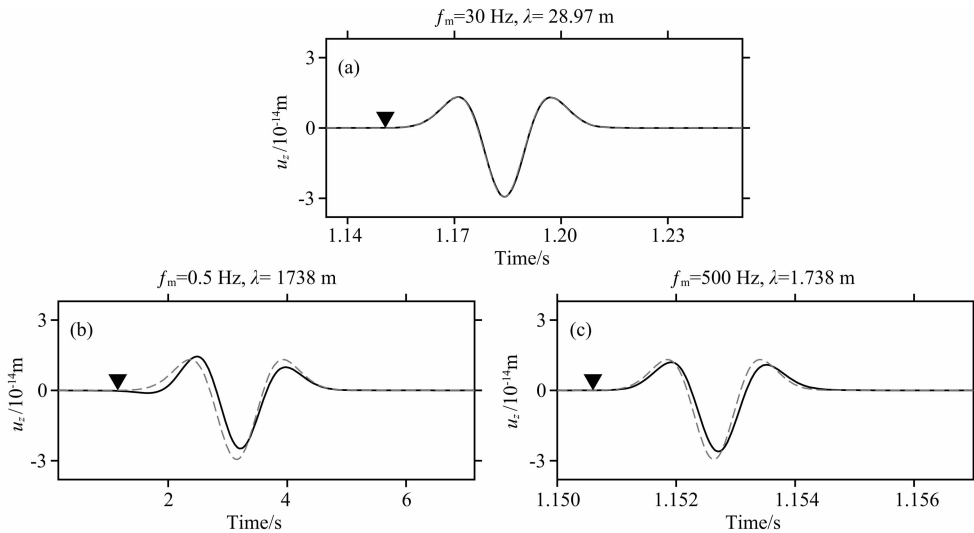


图 7 轴向应力源激发波场的半解析解(实线)和解析解(虚线)的对比

图(a—c)分别是 Ricker 子波主频为 30, 0.5, 500 Hz 时检波器的垂向分量. 倒三角形是 SV 波的理论到时.

Fig. 7 Comparison of the semi-analytical solution (solid line) and the analytical solution (dashed line) of wave field excited by a axial stress source

Panels (a)—(c) are vertical components of seismograms with the peak frequency of Ricker wavelet being 30, 0.5, 500 Hz respectively. The inverted triangle denotes the theoretical arrival time of SV waves.

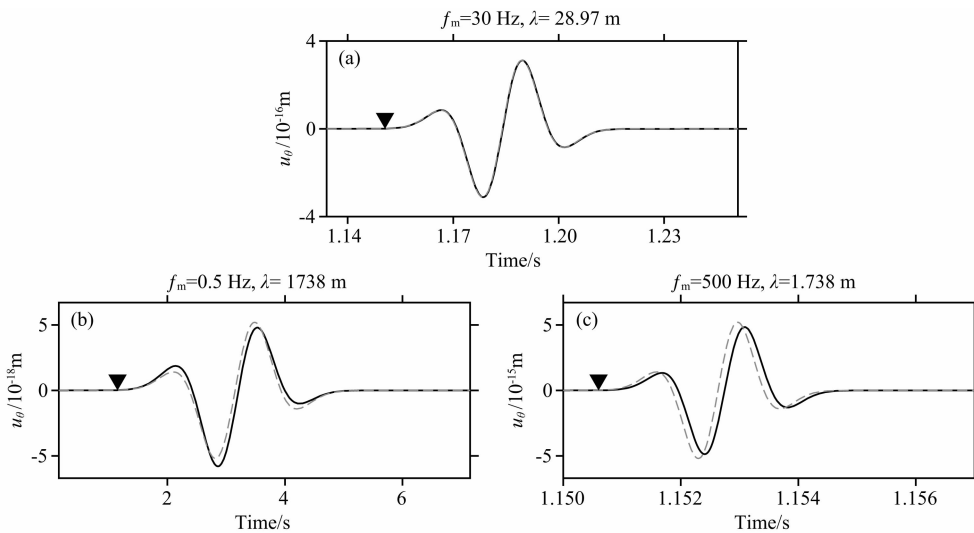


图 8 旋转应力源激发波场半解析解(实线)和解析解(虚线)的对比

图(a—c)分别是 Ricker 子波主频为 30, 0.5, 500 Hz 时检波器的横向分量. 倒三角形是 SH 波的理论到时.

Fig. 8 Comparison of the semi-analytical solution (solid line) and the analytical solution (dashed line) of wave field excited by a torsional stress source

Panels (a)—(c) are transverse components of seismograms with the peak frequency of Ricker wavelet being 30, 0.5, 500 Hz respectively. The inverted triangle denotes the theoretical arrival time of SH waves.

振荡器震源、随钻震源, 或者与径向震源来共同模拟一些较为复杂的震源. 计算结果如图 7 所示. 因为轴向震源在检波器处的位移场只有垂向分量, 所以图中对比的是解析解和数值解的垂向分量. 轴向震源的远场波形是 Ricker 子波(图 7), 因此两者绝对振幅都无明显变化. 当满足“小井孔”假设和“远场”假设时, 两者差异很小(图 7a), 而当“远场”假设失效

(图 7b)或“小井孔”假设失效(图 7c)时, 两者之间会有较为明显的差异.

4.4 旋转应力震源

旋转应力震源可能由钻头旋转产生, 也有可能由钻缆与井壁的摩擦产生. 通常情况下旋转应力源的强度较小, 产生的地震波能量也较弱. 计算结果如图 8 所示. 与前两种震源产生 P-SV 波不同, 旋转应

力震源只产生 SH 波,所以图 8 中比较的是位移的横向分量.由于旋转震源远场波形是 Ricker 子波的导数,所以两者的绝对振幅与频率呈现正相关.在“远场”假设(图 8b)或“小井孔”假设(图 8c)失效时,两者同样会出现明显的差异.

5 结论

本文采用沿最速下降路径的数值积分计算了井中震源的远场波场.由于沿该路径的积分函数不存在任何振荡,因此获得了高精度的数值积分.

最速下降积分在计算时可以将 P 波和 SV 波分离,避免了波形之间的互相干扰.该方法保留了对波数域极点和支线进行分析的可能性,可以解析地得到面波,折射波出现的位置和波形.模型试验表明,当“小井孔”和“远场”假设不满足时,近似条件下解析解的振幅和波形相对于半解析解都会有明显的偏差,使用半解析解能获得更准确的波场信息.

附录 A 井中震源的频率波数域解

柱坐标系下的波动方程为

$$\begin{aligned} & (\lambda + 2\mu) \left(\frac{\partial^2 u_r}{\partial r^2} + \frac{1}{r} \frac{\partial u_r}{\partial r} \right) + \frac{\mu}{r^2} \frac{\partial^2 u_r}{\partial \varphi^2} + \mu \frac{\partial^2 u_r}{\partial z^2} - \frac{(\lambda + 2\mu)}{r^2} u_r + \frac{1}{r} (\lambda + \mu) \frac{\partial^2 u_\theta}{\partial r \partial \theta} - \frac{1}{r^2} (\lambda + 3\mu) \frac{\partial u_\theta}{\partial \theta} \\ & + (\lambda + \mu) \frac{\partial^2 u_z}{\partial r \partial z} = \rho \frac{\partial^2 u_r}{\partial t^2}, \\ & \frac{1}{r} (\lambda + \mu) \frac{\partial^2 u_r}{\partial r \partial \theta} + \frac{1}{r^2} (\lambda + 3\mu) \frac{\partial u_r}{\partial \varphi} + \mu \left(\frac{\partial^2 u_\theta}{\partial r^2} + \frac{1}{r} \frac{\partial u_\theta}{\partial r} \right) + \frac{\lambda + 2\mu}{r^2} \frac{\partial^2 u_\theta}{\partial \theta^2} + \mu \frac{\partial^2 u_\theta}{\partial z^2} - \frac{\mu}{r^2} u_\theta \\ & + \frac{1}{r} (\lambda + \mu) \frac{\partial^2 u_z}{\partial \theta \partial z} = \rho \frac{\partial^2 u_\theta}{\partial t^2}, \\ & (\lambda + \mu) \left(\frac{\partial^2 u_r}{\partial r \partial z} + \frac{1}{r} \frac{\partial u_r}{\partial z} \right) + \frac{1}{r} (\lambda + \mu) \frac{\partial^2 u_\theta}{\partial \theta \partial z} + \mu \left(\frac{\partial^2 u_z}{\partial r^2} + \frac{1}{r} \frac{\partial u_z}{\partial r} + \frac{1}{r^2} \frac{\partial^2 u_z}{\partial \varphi^2} \right) + \lambda \frac{\partial^2 u_z}{\partial z^2} = \rho \frac{\partial^2 u_z}{\partial t^2}. \end{aligned} \quad (A1)$$

将势函数表达式 $\mathbf{u} = \nabla \phi + \nabla \times [\nabla \chi \mathbf{e}_z + \nabla \times (\psi \mathbf{e}_z)]$ 代入上式可得

$$\left(\nabla^2 - \frac{\partial^2}{\partial z^2} \right) \left[\mu \nabla^2 \chi - \rho \frac{\partial^2 \chi}{\partial t^2} \right] = 0, \quad (A2)$$

其中 $\nabla^2 = \frac{\partial^2}{\partial r^2} + \frac{1}{r} \frac{\partial}{\partial r} + \frac{1}{r^2} \frac{\partial^2}{\partial \theta^2} + \frac{\partial^2}{\partial z^2}$. 这与各向同性介质中 SH 波与 P, SV 波解耦的现象一致. 剩下的两个关于 P-SV 的方程为

$$\begin{aligned} & \left(\nabla^2 - \frac{\partial^2}{\partial z^2} \right) \left\{ \left[(\lambda + 2\mu) \nabla^2 \phi - \rho \frac{\partial^2 \phi}{\partial t^2} \right] + \frac{\partial}{\partial z} \left[\mu \nabla^2 \psi - \rho \frac{\partial^2 \psi}{\partial t^2} \right] \right\} = 0, \\ & \frac{\partial}{\partial z} \left[(\lambda + 2\mu) \nabla^2 \phi - \rho \frac{\partial^2 \phi}{\partial t^2} \right] + \left(\frac{\partial^2}{\partial z^2} - \nabla^2 \right) \left[\mu \nabla^2 \psi - \rho \frac{\partial^2 \psi}{\partial t^2} \right] = 0. \end{aligned} \quad (A3)$$

在柱坐标系下,通常用柱面波分解代替平面波分解来求解波动方程.柱面波可以表示为 $H_n^{(1)}(qr) e^{-i\omega t}$, 其中 $H_n^{(1)}(z)$ 是 Bessel 函数的一种,也叫做第一类 Hankel 函数, q 是径向波数.在前人的研究工作中(Abo-Zena, 1977; Blair, 2007),有些研究者采用修正 Bessel 函数,但是 $e^{-i\omega t}$ 和 $H_n^{(1)}(z)$ 的组合可以更直观描述向外传播的柱面波.设

$$\begin{aligned} \phi(r, \theta, z, t) &= \left(\frac{1}{2\pi} \right)^2 \sum_{n=-\infty}^{\infty} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \tilde{\Phi}_n(r, h, \omega) \exp[i(hz + n\theta - \omega t)] dh d\omega, \\ \psi(r, \theta, z, t) &= \left(\frac{1}{2\pi} \right)^2 \sum_{n=-\infty}^{\infty} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \tilde{\Psi}_n(r, h, \omega) \exp[i(hz + n\theta - \omega t)] dh d\omega, \\ \chi(r, \theta, z, t) &= \left(\frac{1}{2\pi} \right)^2 \sum_{n=-\infty}^{\infty} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \tilde{X}_n(r, h, \omega) \exp[i(hz + n\theta - \omega t)] dh d\omega, \end{aligned} \quad (A4)$$

其中

$$\tilde{\Phi}_n(r, h, \omega) = f_1 H_n^{(1)}(qr), \tilde{\Psi}_n(r, h, \omega) = f_2 H_n^{(1)}(qr), \tilde{X}_n(r, h, \omega) = f_3 H_n^{(1)}(qr) \quad (A5)$$

是柱面波,也可以看作是频率波数域的解.将(A4)、(A5)代入(A3)式中,并考虑到 Hankel 函数的定义

$$\frac{d^2}{dr^2} H_n^{(1)}(qr) + \frac{1}{r} \frac{d}{dr} H_n^{(1)}(qr) + \left(q^2 - \frac{n^2}{r^2} \right) H_n^{(1)}(qr) = 0, \quad (A6)$$

可以得到

$$\begin{aligned} q^2 \{ [(\lambda + 2\mu)(q^2 + h^2) - \rho\omega^2] f_1 + ih[\mu(q^2 + h^2) - \rho\omega^2] f_2 \} H_n^{(1)}(qr) &= 0, \\ \{ ih[-(\lambda + 2\mu)(q^2 + h^2) + \rho\omega^2] f_1 - q^2[\mu(q^2 + h^2) - \rho\omega^2] f_2 \} H_n^{(1)}(qr) &= 0, \\ q^2 [\mu(q^2 + h^2) - \rho\omega^2] f_3 H_n^{(1)}(qr) &= 0. \end{aligned} \quad (A7)$$

(A7)式是关于 f_1, f_2, f_3 的线性方程组, 如要得到非零解, 必须满足

$$q^4 (q^2 + h^2) [(\lambda + 2\mu)(q^2 + h^2) - \rho\omega^2] [\mu(q^2 + h^2) - \rho\omega^2] [\mu(q^2 + h^2) - \rho\omega^2] (H_n^{(1)}(qr))^3 = 0, \quad (A8)$$

其中当 $q = 0$ 时, $H_n^{(1)}(qr)$ 是一个奇点, 上述的推导过程不成立, 所以现在只考虑 $q \neq 0$ 的情况. 这时, (A8) 式可以简化成

$$[(\lambda + 2\mu)(q^2 + h^2) - \rho\omega^2] [\mu(q^2 + h^2) - \rho\omega^2] [\mu(q^2 + h^2) - \rho\omega^2] = 0, \quad (A9)$$

可以看到, P, SV, SH 波在各向同性介质中都是解耦的. 从 (A9) 式可以得到 q 的三对共轭解, 根据 Sommerfeld 辐射条件, 我们选择 $\text{Im}q > 0$ 的三个解来保证波场解在无穷远处的能量有限. 当 $\text{Im}q = 0$ 时, 我们选择 $\omega \text{Re}q > 0$ 的解来保证波场是向外传播的. 所以, 势函数在频率波数域的解为

$$\tilde{\Phi}_n(r, h, \omega) = f_1 H_n^{(1)}(q_1 r), \tilde{\Psi}_n(r, h, \omega) = f_2 H_n^{(1)}(q_2 r), \tilde{X}_n(r, h, \omega) = f_3 H_n^{(1)}(q_3 r), \quad (A10)$$

其中 $q_1 = [\rho\omega^2/(\lambda + 2\mu) - h^2]^{1/2}$, $q_2 = q_3 = [\rho\omega^2/\mu - h^2]^{1/2}$. $f_{1,2,3}$ 是三个待定的参数, 可以通过边界条件来确定. 设边界条件为

$$\begin{aligned} \sigma_{rr} |_{r=a} &= b_1(\theta, z, t), \\ \sigma_{r\theta} |_{r=a} &= b_2(\theta, z, t), \\ \sigma_{rz} |_{r=a} &= b_3(\theta, z, t). \end{aligned} \quad (A11)$$

将边界条件也变换到频率波数域, 可以得到关于 $f_{1,2,3}$ 的线性方程组

$$\mathbf{D}\mathbf{f} = \tilde{\mathbf{B}}, \quad (A12)$$

其中 $\tilde{\mathbf{B}} = [\tilde{B}_{1n}, \tilde{B}_{2n}, \tilde{B}_{3n}]^T$ 是 $b_{1,2,3}(\theta, z, t)$ 在频率波数域中表示,

$$\tilde{B}_{jn} = \frac{1}{2\pi} \int_{-\pi}^{\pi} \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} b_j(\theta, z, t) \exp[i(-n\theta - hz + \omega t)] dt dz d\theta, \quad (A13)$$

$\mathbf{f} = [f_1, f_2, f_3]^T$, $\mathbf{D} = [D_{ij}]$. D_{ij} 的具体表达式为

$$\begin{aligned} D_{11} &= - \left((\lambda + 2\mu)q_1^2 - 2\mu \frac{n(n-1)}{a^2} + \lambda h^2 \right) H_n^{(1)}(q_1 a) + 2 \frac{1}{a} \mu q_1 H_{n+1}^{(1)}(q_1 a), \\ D_{12} &= - ih \left(2\mu q_2^2 - 2\mu \frac{n(n-1)}{a^2} \right) H_n^{(1)}(q_2 a) + 2 \frac{1}{r} i \mu h q_2 H_{n+1}^{(1)}(q_2 a), \\ D_{13} &= 2in\mu \left(\frac{n-1}{a^2} H_n^{(1)}(q_3 a) - q_3 \frac{1}{a} H_{n+1}^{(1)}(q_3 a) \right), \\ D_{21} &= 2i\mu \left(\frac{n(n-1)}{a^2} H_n^{(1)}(q_1 a) - q_1 \frac{n}{a} H_{n+1}^{(1)}(q_1 a) \right), \\ D_{22} &= - 2\mu h \left(\frac{n(n-1)}{a^2} H_n^{(1)}(q_2 a) - q_2 \frac{n}{a} H_{n+1}^{(1)}(q_2 a) \right), \\ D_{23} &= \mu \left(\left(q_3^2 - \frac{2n(n-1)}{a^2} \right) H_n^{(1)}(q_3 a) - q_3 \frac{2}{a} H_{n+1}^{(1)}(q_3 a) \right), \\ D_{31} &= 2i\mu h \left[\frac{n}{a} H_n^{(1)}(q_1 a) - q_1 H_{n+1}^{(1)}(q_2 a) \right], \\ D_{32} &= \mu (q_2^2 - h^2) \left[\frac{n}{a} H_n^{(1)}(q_1 a) - q_2 H_{n+1}^{(1)}(q_2 a) \right], \\ D_{33} &= - \frac{1}{a} n \mu h q_3 H_n^{(1)}(q_3 a). \end{aligned} \quad (A14)$$

从 (A12)、(A14) 可以看出, 任意类型的井中震源都可以产生 P, SV 和 SH 波.

轴对称情况对应于 $n=0$ 的情况. 这时 D_{ij} 的表达式退化

$$\begin{aligned}
D_{11} &= -((\lambda + 2\mu)q_1^2 + \lambda h^2)H_0^{(1)}(q_1 a) + 2\frac{1}{a}\mu q_1 H_1^{(1)}(q_1 a), \\
D_{12} &= -2i\mu h q_2^2 H_0^{(1)}(q_2 a) + 2\frac{1}{r}i\mu h q_2 H_1^{(1)}(q_2 a), \\
D_{23} &= \mu q_3 \left(q_3 H_0^{(1)}(q_3 a) - \frac{2}{a}H_1(q_3 a) \right), \\
D_{31} &= -2i\mu h q_1 H_1^{(1)}(q_2 a), \\
D_{32} &= -\mu q_2 (q_2^2 - h^2)H_1^{(1)}(q_2 a), \\
D_{13} &= D_{21} = D_{22} = D_{33} = 0.
\end{aligned} \tag{A15}$$

解这个方程组得到 $f_{1,2,3}$ 后,将(A10)代入势函数的表达式,得到频率波数域的位移解表达式:

$$\begin{bmatrix} \tilde{U}_r \\ \tilde{U}_\theta \\ \tilde{U}_z \end{bmatrix} = \begin{bmatrix} -q_1 H_1^{(1)}(q_1 r) & -ihq_2 H_1^{(1)}(q_2 r) & 0 \\ 0 & 0 & q_3 H_1^{(1)}(q_3 r) \\ ihH_0^{(1)}(q_1 r) & q_2^2 H_0^{(1)}(q_2 r) & 0 \end{bmatrix} \begin{bmatrix} f_1 \\ f_2 \\ f_3 \end{bmatrix}, \tag{A16}$$

记(A16)式中的矩阵为 \mathbf{K} ,则

$$\tilde{\mathbf{U}} = \mathbf{K} \frac{\text{adj} \mathbf{D} \tilde{\mathbf{B}}}{\det \mathbf{D}} \tag{A17}$$

其中 $\det \mathbf{D}$ 是 \mathbf{D} 的行列式, $\text{adj} \mathbf{D}$ 是 \mathbf{D} 的伴随矩阵.

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